

# Equatorial convergence of India and changes in atmospheric CO<sub>2</sub> and climate during the Cenozoic

D.V. KENT<sup>1,2\*</sup>, G. MUTTONI<sup>3</sup>

<sup>1</sup> Earth & Planetary Sciences, Rutgers University, Piscataway, NJ 08854, USA (\*correspondence: dvk@rutgers.edu)

<sup>2</sup> Lamont-Doherty Earth Observatory, Palisades, NY 10964, USA

<sup>3</sup> Dipartimento di Scienze della Terra, Università di Milano, 20133 Milano, Italia (giovanni.muttoni1@unimi.it)

Conflicting estimates of seafloor production rates over the Late Cretaceous and Cenozoic have variously indicated generally decreasing (e.g., (1, 2)), steady (3), or even increasing trends (4) that raise questions about the often presumed linkages of variations in sea-floor spreading with changes in eustatic sea level (e.g., (5)) and especially volcanic outgassing as a principal control on the evolution of atmospheric CO<sub>2</sub> (6). Elaborating on ideas expressed several years ago by D. Schrag (7) and J. Edmond and Y. Huh (8), we suggest that India's northward flight and collision with Asia was a major driver of atmospheric CO<sub>2</sub> concentration (*p*CO<sub>2</sub>) and thus global climate in the late Cretaceous and Cenozoic (9). Subduction of Tethyan oceanic crust with a carpet of carbonate-rich pelagic sediments deposited during transit beneath the high productivity equatorial belt resulted in a component flux of CO<sub>2</sub> delivery to the atmosphere that maintained high *p*CO<sub>2</sub> levels and warm climate until the decarbonation factory waned with the collision of Greater India with Asia at ~50 Ma (10), closely coinciding with the Early Eocene climatic optimum. At about this time, the India continent and the highly weatherable Deccan Traps drifted into the equatorial humid belt where uptake of CO<sub>2</sub> by silicate weathering further perturbed the equilibrium towards progressively lower *p*CO<sub>2</sub> levels (11) and a cooling trend that eventually triggered the expansion of Antarctic ice sheets in the earliest Oligocene (12), even if global seafloor production rates remained steady.

[1] Engebretson et al. (1992) *GSA Today* **2**, 93.

[2] Muller et al. (2008) *Science* **319**, 1357.

[3] Rowley (2002) *Geological Society of America Bulletin* **114**, 927.

[4] Cogné & Humler, (2006) *Geochemistry, Geophysics, Geosystems* **7**, Q03011.

[5] Gaffin (1987) *American Journal of Science* **287**, 596.

[6] Berner (1994) *American Journal of Science* **294**.

[7] Schrag (2002) *Geochimica and Cosmochimica Acta* **66**, A688.

[8] Edmond & Huh (2003) *Earth and Planetary Science Letters* **216**, 125.

[9] Kent & Muttoni (2008) *Proceedings of the National Academy of Science* **105**, 16065.

[10] Garzanti et al. (1987) *Geodinamica Acta* **1**, 297.

[11] Pagani et al. (2005) *Science* **309**, 600.

[12] DeConto & Pollard (2003) *Nature* **421**, 245.